



# **A simple model to estimate emission of wind-blown particles from a granular bed in comparison to wind tunnel experiments**

M.C.S. Ferreira, B. Furieri, Ahmed Ould El Moctar, J.-L. Harion, Alexandre Valance, P. Dupont, N.C. Reis, J.M. Santos

## **► To cite this version:**

M.C.S. Ferreira, B. Furieri, Ahmed Ould El Moctar, J.-L. Harion, Alexandre Valance, et al.. A simple model to estimate emission of wind-blown particles from a granular bed in comparison to wind tunnel experiments. *Geomorphology*, 2019, 335, pp.1-13. 10.1016/j.geomorph.2019.03.004 . hal-02089227

**HAL Id: hal-02089227**

**<https://univ-rennes.hal.science/hal-02089227>**

Submitted on 6 Jun 2019

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Geomorphology

# A simple model to estimate emission of wind-blown particles from a granular bed in comparison to wind tunnel experiments

M. C. S. Ferreira, B. Furieri, A. Ould El Moctar, J.-L. Harion, A. Valance, P. Dupont, N. C. Reis Jr, J. M. Santos

*Department of Environmental Engineering, Universidade Federal do Espírito Santo, 29075-910, Vitória, ES, Brazil*

*Department of Mathematics, Instituto Federal do Espírito Santo, 29040-780, Vitória, ES, Brazil*

*Laboratoire de Thermique et Energie de Nantes, UMR CNRS 6607, 44306 Nantes, France  
IMT Lille-Douai, 59500 Douai, France*

*Institut de Physique de Rennes, CNRS UMR 6251, Université de Rennes, 35042 Rennes Cedex, France*

*LGCGM, EA3913, Université de Rennes, 35042 Rennes Cedex, France*

## Abstract

Dust emissions due to aeolian erosion of exposed granular materials are strongly influenced by grain size distribution. Non-erodible particles that are too heavy to be lifted into the air play a protective role in the aeolian erosion process attenuating emission, which is known as the pavement phenomenon. To date, there is no approach that reliably predicts the reduction in emissions caused by their presence on an aggregate surface. In this work, an analytical model was developed to quantify emissions from particle beds with a wide size distribution. As non-erodible particles accumulate, changes in surface characteristics create increasing shelter for the erodible portion of the bed until the shear on the erodible surface reaches a minimum and emissions cease. The proposed emission model describes the relationship between this minimum value of wind shear and the eroded depth of the bed after the pavement, which in turn gives the emitted mass. In addition, wind tunnel experiments were carried out in order to broaden knowledge of the pavement phenomenon and validate the modelling. A bimodal particle size distribution of sand with erodible and non-erodible particles was used for the tested velocities. Three experimental measurements were carried out: (i) continuous weighing of the emitted mass, (ii) eroded depth of the bed at regular time intervals and (iii) final cover rates of the non-erodible particles using digital analysis of sand bed pictures after experiments. Good agreement between the modelling and experimental results was found. The emission model proposed herein is a simple algebraic expression that demands low computational effort. This approach may serve as a base for an emission model for application in granular materials stockpiles.

**Keywords:** Wind Erosion, Non-erodible Particles, Fugitive emissions, Emission model

## 1. Introduction

Wind erosion is a natural process characterised by particle entrainment, transport and deposition due to the action of wind. The understanding of the phenomenon is important in several research fields such as land degradation in agricultural areas [58,2], deserts expansion [20], dunes morphology and dynamics [6,9], dust storms [29], design of porous

fences to prevent soil erosion [31] and air pollution [5]. Particulate air pollutants have diverse chemical compositions and cover a wide range of sizes, from  $0.001\text{ nm}$  to  $100\text{ }\mu\text{m}$ . Although the particles larger than  $10\text{ }\mu\text{m}$  have a shorter atmospheric lifetime and settle closer to the sources of emission, they can cause annoyance related to dirtiness, eye irritation, cough and allergic reaction. The long-range transport of the smallest particles, such as  $PM_{10}$  ( $< 10\text{ }\mu\text{m}$ ),  $PM_{2.5}$  ( $< 2.5\text{ }\mu\text{m}$ ) and ultrafine particle fractions ( $< 0.1\text{ }\mu\text{m}$ ), can reach atmospheric lifetimes of days, weeks, or even years [46]. These particles are responsible for a variety of serious harmful effects on health, especially related to cardiovascular and respiratory diseases [55].

Emissions due to aeolian erosion of exposed granular materials are strongly influenced by the grain size distribution possibly presenting coarse particles that are not lifted by wind flow [17,48,49,57]. In the range of investigated wind velocities, the basal shear stress is insufficient to lift the large particles of  $1\text{ mm}$  in diameter, referred hereafter as non-erodible particles. Wind erosion processes and emission estimation have been studied by numerical and experimental approaches. Several experimental studies have shown that the presence of these non-erodible particles affects dust emission limiting the availability of particles supply and promoting a temporal decrease in emitted mass flux [33,30,32,11,53,25]. As the surface is eroded, the coarse particles accumulate and protect the bed surface from erosion, which is known as the pavement phenomenon. Therefore, it is important to consider temporal variability in surface conditions to develop wind erosion models [41,54,37]. Nonetheless, aeolian transport is often viewed as a steady-state process, and most available emission models are applied to particle beds with a homogeneous size distribution. Some models account for the wide range of particle sizes in the bed [35,1,28], however, modifications of soil size distribution with time are not usually considered. The work of [12] was the first attempt to fill this gap. The authors developed a stochastic model that quantifies the temporal evolution of the emitted mass flux in a bed of granular material with a wide size distribution exposed to a turbulent flow. To model pavement, [12] assumed that erosion is finalised when the bed is completely overlaid by non-erodible particles. However, experimental results showed that when erosion stops, potential erodible particles still remain on the surface [11,17]. The accurate prediction of the phenomenon is a complex task due to the uncertainty in the transport processes. [42] describe in their paper this inherent uncertainty in the saltation phenomenon. [36] presented a new methodology for characterizing high-frequency fluctuations of aeolian saltation flux.

The present work aims to investigate the pavement phenomenon and propose an emission model, which includes a new approach to characterise the influence of non-erodible particles accumulation. The non-erodible particles protect the erodible surface by both covering the surface and creating downstream wake zones of reduced wind shear stress on the intervening bare surface [19,23,44,38,40]. As the concentration of larger particles in the bed increases, the mean shear on erodible surfaces decreases until it is no longer significant to cause emissions. Thus, pavement modelling in the proposed model is related to the wall shear stress evolution as wind erosion modifies the surface. The partition of shear between non-erodible particles and adjacent erodible ones is investigated using numerical simulations of turbulent flow over a rough surface and a mathematical formulation to associate the geometrical characteristics of the non-erodible particles and the mean friction velocity on the erodible surface (Section 2). Based on this formulation, the final conditions of the bed, such as the eroded depth and the cover rate of non-erodible particles, were determined by assuming that after pavement, the friction velocity reaches a minimum value. Then, besides providing the emitted mass from a particle bed, the model gives the final state of the bed topography (Section 3). The primary objective of this study is to propose a simple model for dust and particle emission in practical situations such as in industrial sites of storage piles of granular material or in semi-arid

environments (soil erosion) from the knowledge of the distribution of the surface friction at the local level. An experimental study in a wind tunnel, described in Section 3, was conducted to validate the modelling results (final characteristics of the bed surface and global emissions estimates).

The primary objective of this study is to propose accurate and simple models for dust and particle emission in practical situations such as in industrial sites of storage piles of granular material or in semi-arid environments (soil erosion). The main ingredient of the model is the knowledge of the distribution of the surface friction at the local level. The only requirement to apply this model at regional scale for dust emission prediction is to assess the surface shear distribution linked to the local roughness. The purpose of the wind tunnel study is to validate this model in well-controlled situations.

## 2. Erosion model

### 1. Grain entrainment and saltation

Wind erosion of a particle bed is initiated by aerodynamic entrainment, whenever the velocity exceeds a critical value called the static threshold. The onset of grain motion is commonly described by the Shields number  $\Theta$ , defined as the ratio between the shear force exerted by a fluid on a particle at the bed surface and the effective particle weight:

$$\Theta = \frac{\rho u^*{}^2}{(\rho_p - \rho)gD}, \quad (1)$$

where  $u^*$  is the friction velocity,  $\rho$  is the fluid density,  $\rho_p$  is the particle density,  $g$  is gravity and  $D$  is the particle diameter.

[3] investigated the static threshold Shields number  $\Theta_s$  (corresponding to the static threshold friction velocity  $u_{ts}^*$ ) based on the balance of forces acting on a particle. The author found that  $\Theta_s$  is nearly constant for large particles ( $\Theta_s \approx 0.01$ ), which implies that  $u_{ts}^*$  increases with  $\sqrt{D}$  in these cases. However, for small particles,  $\Theta_s$  is not constant and increases rapidly as diameter decreases. As particle size decreases, inter-particle cohesion forces can no longer be neglected and its effect plays a dominant role in controlling threshold motion conditions. Since the pioneering work of [3], other studies have been devoted to investigating the threshold Shields number and erodibility conditions for the particles [24,15,8,14,13]. In the present work, the erodibility of the particles was assessed by a take-off criterion obtained by [47], who presented a simple expression relating the threshold friction velocity  $u_{ts}^*$  and the particle diameter  $D$ :

$$u_{ts}^* = 0.11 \sqrt{\frac{\rho_p - \rho}{\rho} gD + \frac{\gamma}{\rho D}}, \quad (2)$$

where  $\gamma$  is a surface energy that characterises the cohesion.

Equation 2 implies that if inter-particle cohesion is considered,  $u_{ts}^*$  is proportional to  $\sqrt{c_1 D + c_2 D^{-1}}$  (instead of  $\sqrt{D}$ ). For large particles, the term  $c_1 D$  dominates over  $c_2 D^{-1}$ , which is consistent with  $\Theta_s \approx 0.01$ . This criterion is in good agreement with a large number of wind tunnel measurements [27]. [47] recommended values of  $\gamma$  ranging between  $1.65 \times 10^{-4}$  and  $5.00 \times 10^{-4} \text{ kg/s}^2$ . In this work,  $\gamma = 2.86 \times 10^{-4} \text{ kg/s}^2$  was used, as obtained by [26] experimentally fitting Equation 2 to the threshold required to lift erodible particles. The take-off criteria (illustrated in Fig. 1) allows the estimation of the range of particle sizes

liable to take-off by aerodynamic entrainment. For a given friction velocity  $u^*$ , there will be a range of critical diameters in which particles are emitted. Outside of this range, particles remain on the surface due to gravitational forces for larger particles and high cohesion forces for smaller particles.

**Figure 1.** Criteria for the definition of particles erodibility obtained by [47]

In the early stage of erosion, aerodynamic forces are mainly responsible for the entrainment of particles. Transport is initiated by aerodynamic entrainment of a small number of particles if the static threshold  $\Theta_s$  is reached. These particles may hop along the surface, and the impact of saltating grains on the bed surface can eject other particles. After saltation has initiated, subsequent lifting of surface particles occurs predominantly due to impact. Then, transport can be sustained even below the static threshold. The minimum velocity at which the combined action of wind forces and saltation impacts are capable of retaining particles in movement is called the dynamic threshold ( $\Theta_D$ ). The dynamic threshold is lower than the static threshold and is also fairly constant ( $\Theta_D \approx 0.008$ ) for large grains [7,22,8,14]. For smaller particles (typically  $\lesssim 100\mu m$ ), there is an increase in  $\Theta_D$  due to the importance of cohesion forces. The expression of the threshold friction velocity versus the particle diameter presented herein (Equation 2) was developed for beds composed of particles with monodispersed size distribution, without roughness elements as non-erodible particles. The effects of the presence of roughness elements are taken into account by the shear stress distribution between the roughness elements and the ground surface.

## 2. Shear stress partition

The shear stress partition has been investigated by several researchers by means of theoretical [43,35] and experimental [34,10] approaches. Despite the importance of these works, they present shortcomings, most of all concerning large variations found in model parameters and the limited range validity for certain types of rough surfaces. Computational Fluid Dynamics (CFD) is a valuable alternative that can be applied to simulate the flow over surfaces with different shapes and arrangements of roughness elements. Different configurations may be too costly and time consuming to be predicted by experimental techniques.

[45] proposed that the total drag  $\tau$  imparted to a rough surface can be written as:

$$\tau = \rho u^{*2} = \tau_R + \tau_S, \quad (3)$$

where  $\tau_R$  is the pressure drag on the roughness elements and  $\tau_S$  is the friction drag on the ground surface.

The shear stress  $\tau_S'$  acting on the exposed surface (erodible), which drives wind erosion, is related to  $\tau_S$  by:

$$\tau_S' = \rho u_s^{*2} = \frac{\tau_S}{1 - CR}, \quad (4)$$

where  $u_s^*$  is the mean friction velocity on the erodible fraction of a rough surface and  $CR$  is the cover rate of the surface by the roughness elements.

A theoretical model was proposed by [43], based on the idea that the wake and drag properties of an isolated roughness element can be characterised by an effective shelter area and volume. The shelter area describe the surface stress deficit behind the element and the shelter volume, the attenuation of drag on other obstacles in the element wake. It was assumed that the combined effective shelter can be calculated by randomly superimposing individual shelters. However, the validity of this hypotheses is limited to low values of roughness densities  $\lambda$ . As

$\lambda$  increases, element wake interactions become stronger and their collective effect cannot be described by superimposition. [43] model is given by:

$$\frac{\tau_s'}{\tau} = \frac{1}{(1 - \sigma\lambda)(1 + \beta'\lambda)}, \quad (5)$$

where  $\beta'$  is the ratio of drag coefficient of an individual roughness element to the drag coefficient of the bare surface. The parameter  $\beta'$  accounts for roughness element shape effects and entirely controls the partition of drag.

The [43] model has been assessed by several works in order to test and refine the parameters  $\beta'$  [56,10,4,52]. A general conclusion is that the proposed formulation is valid, nonetheless, presents shortcomings. [39] and [10] found that the parameter  $\beta'$  has some degree of dependency on the aspect ratio of the roughness elements. The range of values obtained for  $\beta'$  is relatively large, which makes difficult the identification of appropriate values for a specific surface with non-erodible elements [52].

[50] performed numerical simulations with different configurations of rough surfaces to depict the pavement process by determining the friction velocity experienced by the erodible fraction of particles as a function of the changes in bed topography, i.e., the increase in number or the height of the roughness elements. The work analysed the parameter  $R_{fric}$  (Equation 6), which reflects the evolution of the friction velocity applied to the exposed surface as erosion occurs:

$$R_{fric} = \frac{u_s^*}{u_0^*}, \quad (6)$$

where  $u_0^*$  is the mean friction velocity on a surface without roughness elements. Then, [50] developed an empirical relationship to include the cover rate and particles geometrical parameters on  $R_{fric}$ :

$$1 - R_{fric} = A(CR)^M \left( \frac{S_{frontal}}{S_{floor}} \right)^N, \quad (7)$$

where  $S_{frontal}$  and  $S_{floor}$  are respectively the frontal and the basal areas of a cylindrical roughness element, and  $A, M$  and  $N$  are coefficients determined by the numerical simulations.

[16] found that Equation 7 is also valid for surfaces containing particles with different aspect ratios (non-uniform distribution of roughness elements). However, this formulation was tested for cover rates lower than 12%. Nonetheless, the particle size distribution found in nature, and particularly in granular materials from storage yards of industrial sites, can present higher cover rate values. Therefore, in the present study, additional numerical simulations were carried out in order to adjust the values of the coefficients  $A, M$  and  $N$  for a larger range of cases. The studied configurations correspond to a bed of granular material, in which the non-erodible particles are represented by non-uniform cylindrical roughness elements randomly distributed, emerging at various levels from the surface (see Fig. 2). The transport of particles is not taken into account in the numerical simulations.

**Figure 2. Random positioning of the non-uniform roughness elements in the numerical simulations**

The parameters that define each test are: the range of roughness elements diameter  $D_r$  and height  $h_r$  (proportionally to  $D_r$ ) and the number of roughness elements  $n_r$  ( $n_r = 0$  corresponds to a surface without roughness elements, i.e.,  $u_s^* = u_0^*$ ). The ratio  $S_{frontal}/S_{floor}$  and the cover rate  $CR$  are determined for each simulated case according to:

$$\frac{S_{frontal}}{S_{floor}} = \frac{4 \bar{h}_r}{\pi \bar{D}_r}, \quad (8)$$

$$CR = \frac{\frac{\pi}{4} \sum_r^{n_r} D_r^2}{S_{bed}}, \quad (9)$$

where  $\bar{D}_r$  and  $\bar{h}_r$  are, respectively, the mean diameter and the mean emergent height of the roughness elements, and  $S_{bed}$  is the area of the bed. Table 1 presents the configurations of the simulated beds.

**Table 1. Configurations of particle bed covered by cylindrical roughness elements tested in the numerical simulations.  $u_S^*$  is the mean value of the friction velocity on the erodible surface, calculated using the numerical results of the shear stress distribution.  $R_{fric}$  is calculated using Equation 6**

Test	$D_r$ (mm)	$\frac{h_r}{D_r}$ (%)	$n_r$	CR	$\frac{S_{frontal}}{S_{floor}}$	$u_S^*$	$R_{fric}$
1	1.0 — 3.0	20-100	48	15.04	0.80	0.2044	0.58
2	1.0 — 3.0	20-100	62	20.21	0.72	0.1990	0.57
3	1.0 — 3.0	20-100	91	25.02	0.76	0.1829	0.52
4	1.0 — 3.0	20-100	102	27.04	0.77	0.1738	0.50
5	1.0 — 3.0	20-100	116	30.01	0.73	0.1762	0.50
6	1.0 — 3.0	20-100	121	32.04	0.79	0.1588	0.45
7	1.0 — 3.0	20-100	156	35.12	0.77	0.1583	0.45
8	1.0 — 3.0	20-100	158	37.07	0.77	0.1490	0.43
9	1.0 — 3.0	20-100	181	40.12	0.74	0.1530	0.44
10	1.0 — 3.0	60-100	62	20.51	1.00	0.1750	0.50
11	1.0 — 3.0	60-100	113	30.35	1.00	0.1630	0.47
12	1.0 — 3.0	60-100	190	40.21	1.01	0.1398	0.40
13	1.0 — 3.0	80-120	63	20.67	1.31	0.1607	0.46
14	1.0 — 3.0	80-120	114	30.40	1.28	0.1472	0.42
15	1.0 — 3.0	80-120	190	40.02	1.27	0.1250	0.36
16	0.5 — 1.5	80-100	168	15.19	1.14	0.1980	0.57
17	0.5 — 1.5	80-100	238	20.12	1.14	0.1840	0.53

18	0.5	—	80-100	30925.10	1.14	0.1721	0.49
	1.5						
19	0.5	—	80-100	41030.19	1.14	0.1619	0.46
	1.5						
20	0.5	—	80-100	50734.06	1.14	0.1510	0.43
	1.5						
21	0.5	—	140-160	16815.11	1.90	0.1610	0.46
	1.5						
22	0.5	—	140-160	23820.04	1.90	0.1601	0.46
	1.5						
23	0.5	—	140-160	30925.03	1.90	0.1398	0.40
	1.5						
24	0.5	—	140-160	41030.13	1.90	0.1332	0.38
	1.5						
25	0.5	—	140-160	50734.12	1.91	0.1281	0.37
	1.5						

The commercial software FLUENT was used to solve the three-dimensional Reynolds averaged equations of mass and momentum. The turbulence effects were accounted for by using the  $k - \omega$  Shear Stress Transport (SST) model. The turbulence model  $k - \epsilon$  SST is a RANS modelling presenting good accuracy and reduced computational effort required. It has been proven [51,18] that  $k - \omega$  SST model present better agreement with experimental and analytical data for the studied configuration, i.e., fluid flow surrounding obstacles than other RANS models such as  $k - \epsilon$ , etc... The computation domain is represented by a parallelepiped ( $30 \times 30$  mm base and 100 mm height). Periodic conditions were applied to inlet and outlet boundaries, which means that the outlet was used to set the inlet profiles in a cyclic way until a fully developed flow was established. The simulations are initialized with a uniform velocity distribution, and the mass flow is fixed for a corresponding mean longitudinal velocity value equal to 8 m/s. The upper and lateral boundaries were defined as symmetry (normal gradients of all variables are set to zero) and the lower limits (roughness elements and ground walls) as smooth walls with no-slip conditions.

The calculation domain was divided into two parts to allow a good mesh refinement near the wall in the spanwise and streamwise directions, with hexahedral and pentahedral elements, and a coarser mesh with quadrilateral elements elsewhere for better computational efficiency. This is necessary since strong velocity gradients and important fluid interactions occur near the wall. The height of the first mesh is  $z^+ < 5$  to ensure no use of wall functions in the turbulence model. Spanwise and streamwise mesh size in wall units ( $x^+$  and  $y^+$ ) present a maximum value of 10 which occurs far from the roughness elements. The smallest element size of the computational mesh is the cylinder circumference divided by about 20. The computational set up (modelling choices including meshing, turbulence closure and domain configurations) was based on previously validated numerical calculations (cf. [50] and [16]). Fig. 3 shows the friction velocity contours on the erodible surface, determined from the shear stress distribution, for two different simulations: Tests 15 and 16. Test 15 presents higher values of the ratio  $S_{frontal}/S_{floor}$  and  $CR$  than Test 16. Therefore, in Test 15, the wakes formed behind the elements are expanded and interact with the neighbouring particles, leading to a drop in the friction velocity on most of the underlying surface (see Fig. 3(a)) and a lower mean value of the friction velocity on the erodible surface ( $u_s^*$ ), as shown in Table 1.

**Figure 3.** Distribution of the parameter  $R_{fric}$  on the erodible surface for Tests (a) 15 and (b) 16

The coefficients  $A$ ,  $M$  and  $N$  in Equation 7 were recalculated based on the numerical results:



$A = 0.188$ ,  $M = 0.313$  and  $N = 0.216$  through the least squares method. The relative errors between the numerical values of  $u_s^*$  and those calculated by Equations 6 and 7 with the new coefficients do not exceed 5%. The proposed formulation was found to be in good agreement with the numerical results ( $R^2 = 0.97$ ) (see Fig.4).

**Figure 4.** Correlation of  $1-R_{fric}$  results between numerical simulations and fitting of coefficients  $A$ ,  $M$  and  $N$

### 3. Roughness evolution and emission prediction

The main purpose of the erosion model is to determine particle mass emission based on the eroded depth of a particle bed for a given velocity, taking advantage of the formulation in Equation 7. The particles were assumed to be cylindrical, with height equal to the diameter of the base. The input variables of the model are wind velocity ( $U_\infty$ ), bed dimensions and the characteristics of the granular material. Given the particle size distribution, the mass fraction of particles  $\alpha_i$  for each size range of particles with representative diameter  $D_i$  is:

$$\alpha_i = \frac{M_i}{M_{tot}} = \frac{V_i}{V_{tot}}, \quad (10)$$

where  $M_i$  and  $V_i$  are, respectively, the mass and the volume of particles from the range  $i$ ,  $M_{tot}$  is the total mass of the particles and  $V_{tot}$  is the total volume occupied by the particles of the bed, calculated as:

$$V_{tot} = \sum_i V_i = \sum_i n_i V_{D_i}, \quad (11)$$

where  $n_i$  is the number of particles encompassed in the range of  $D_i$  and  $V_{D_i}$  is the volume of one particle with diameter  $D_i$  (for cylindrical particles  $V_{D_i} = \pi D_i^3 / 4$ ).

The total volume of particles is related to the volume of the ( $V_{bed}$ ) through the volume fraction  $\varphi$ :

$$\varphi = \frac{V_{tot}}{V_{bed}}, \quad (12)$$

which can be written as a sum of partial volume fractions  $\varphi_i$  of each diameter  $D_i$ :

$$\varphi = \sum_i \varphi_i, \quad (13)$$

where:

$$\varphi_i = \frac{V_i}{V_{bed}} = \alpha_i \varphi. \quad (14)$$

A spatial uniform distribution of particles was assumed along the height of the bed. For each representative diameter  $D_i$  of a given size range, the number of layers in the bed height was calculated by dividing the height of the bed ( $H_{bed}$ ) by  $D_i$  (height of the particle). Therefore, the number of particles per layer  $n_{i/l}$  for each diameter  $D_i$  is:

$$n_{i/l} = n_i \frac{D_i}{H_{bed}} = \frac{\varphi_i S_{bed} D_i}{V_{D_i}}. \quad (15)$$

Once the number of particles per layer for each representative diameter is known, it is possible to establish a correspondence between the cover rate of non-erodible particles and the eroded depth. The initial cover rate of non-erodible particles ( $hitCR_i$ ) is equivalent to the partial volume fraction  $\varphi_{NE}$  of the non-erodible diameters. Thus,  $CR_i$  can be determined from:

$$CR_i = \alpha_{NE} \varphi, \quad (16)$$

where  $\alpha_{NE}$  is the mass fraction of the non-erodible particles.

Fig. 5 shows the evolution of  $CR$  as a function of  $H$  for a unimodal size distribution of non-erodible particles. In this case, the relation between  $CR$  and  $H$  is:

$$CR = aH + CR_i, \quad (17)$$

where  $a = CR_i/D_{NE}$  and  $D_{NE}$  is the diameter of the non-erodible particles. Equation 17 is valid for the velocities in which the erodibility of the particles does not change and for a monodisperse size distribution of non-erodible particles. If the size distribution of non-erodible particles is polydispersed, the surface covered by non-erodible particles for a given bed depth is calculated by successive summation on all non-erodible particle sizes, and the value of  $a$  is found through curve fitting to these data.

**Figure 5.** Evolution of the surface proportion occupied by non-erodible particles ( $CR$ ) as a function of the depth ( $H$ )

Subsequently, in order to relate airflow velocity and eroded depth, Equation 7 was explored. The progressive emergence of non-erodible roughness elements during erosion was simulated by the successive numerical computations. The emergent height of the exposed non-erodible particles ( $\bar{h}_r$ ) is required for the calculation of  $S_{frontal}$ , which was assumed to be equivalent to the eroded depth of the bed  $H$ . Therefore, using Equation 17, Equation 7 can be rewritten as:

$$1 - R_{fric} = A(aH + CR_i)^M \left( \frac{4H}{\pi D_{NE}} \right)^N. \quad (18)$$

Equation 18 provides a relation between the eroded depth  $H$  and  $R_{fric}$ . As non-erodible particles accumulate, friction velocity on the erodible fraction of the surface  $u_s^*$  decreases until a minimum value in which no take-offs occur; the pavement phenomenon occurs and erosion is terminated. Since emissions due to the impact of particles from saltation are accounted for,  $u_{MIN}^*$  is given by the dynamic threshold of the erodible particles. As cohesion effects werenot reached in the experiments, the dynamic threshold is approximately constant ( $\Theta_D \approx 0.008$ ). Therefore,  $u_{MIN}^* = u_{td}^*(200\mu m) \approx 0.186m/s$ , calculated by Equation 1.

At the point where no erosion occurs, the parameter  $R_{fric}$  reaches a minimum value, that is velocity dependent (given by  $R_{MIN} = u_{MIN}^*/u_0^*(U_\infty)$ ), and the eroded depth reaches a maximum  $H_f$ , which gives Equation 19:

$$1 - R_{MIN} = A(aH_f + CR_i)^M \left( \frac{4H_f}{\pi D_{NE}} \right)^N. \quad (19)$$

In conclusion, given the particle size distribution of the bed and the air flow velocity, it is possible to calculate the final eroded depth of the bed after the pavement phenomenon  $H_f$  (Equation 19). As a result, the emitted mass ( $E_f$ ) can be determined from the emitted volume using a simple algebraic expression, as shown in Equation 20:

$$E_f = (1 - \alpha_{NE})\rho_{bed}H_fS_{bed}, \quad (20)$$

where  $\rho_{bed} = \varphi\rho_p$ .

A synthesis of the model is shown in Fig. 6.

**Figure 6.** Chart of the model

### 3. Experimental study

Wind tunnel experiments were conducted in the present work to investigate the temporal evolution of the pavement process and measure the emitted mass in order to validate the proposed model. Therefore, (i) successive weighing of particle emissions and (ii) measurements of the eroded depth of the bed at regular time intervals were performed and (iii) photographs of the sand bed were taken after the experiments.

The experiments were carried out in a 6.6 m long wind tunnel with a square cross-section (see Fig. 7) coupled to a centrifugal fan controlled by a frequency speed variator. Downstream from the fan, there is a longsteel wind tunnel section. At the beginning of this section, a series of spires was installed to hasten the development of the boundary layer, followed by the test section, in which the glass walls allow measurements and photography. After the wind tunnel test section, there is a trap box where the sand can be collected and weighed. More details of the wind tunnel and its characterisation can be found in [21].

**Figure 7.** (a) Scheme of experimental facilities, composed of: (b) Centrifugal fan, (c) Wind tunnel, (d) Test section and sand trap

Two types of grains with  $\rho_p = 2650 \text{ kg/m}^3$  were used in the experiments: fine sand and coarse sand, with mean diameters of 200 and  $1000\mu\text{m}$ , respectively. The fine sand and the coarse sand are erodible and non-erodible for the three tested velocities: 6.7, 8.5 and 9.6 m/s. Then, in the experiments scenario the smaller grains are larger than the size below which cohesion is important. Therefore, once the cohesive regime is not reached, it can be assumed that the dynamic threshold is constant ( $\Theta_D \approx 0.008$ ).

The experiments comprise wind tunnel tests containing a bimodal granulometry with two different mass fractions of non-erodible particles ( $\alpha_{NE}$ ): 10 and 20%. Fine sand was white and the coarse sand was black to allow for the visualisation of non-erodible particle accumulation. It was therefore possible to calculate the proportion of the surface occupied by non-erodible particles after the pavement phenomenon by means of digital analysis.

The wind tunnel was covered by the bimodal sand bed with  $\rho_{bed} \approx 1600 \text{ kg/m}^3$  ( $\varphi = 0.6$ ), and the fan was switched on for a period of 10 minutes. Nonetheless, the flow was periodically blocked (after 1,2,3,4,5 and 7 minutes) for sampling and weighing of emitted particles in the sand trap and at the exit of the wind tunnel. By weighing this sand, global emissions and the emission rate were estimated. It is important to recall that the transport is in an unsteady state and thus the streamwise sand transport changes with distance and time.

A high quality camera was installed over the wind tunnel top wall (transparent one). Then, the eroded depth of the bed was determined at regular time intervals using a Laser sheet. The aim was to determine the eroded surface depressions in accordance with the displacements of the laser trace. A laser sheet was pointed along the central plane direction of the sand bed. The camera registered the evolution of the eroded surface by photographing the bed every 15 seconds until the end of the 10 minutes. The camera position remained unchanged during the entire experiment. Therefore, it was possible to scale the displacement of the laser with the

number of pixels in the pictures through calibration with the displacement of an object with known size. The final laser images were post-processed using the commercial software Matlab. The emission rate, defined as the emitted mass per unit width per unit time, can be deduced from the evolution of the eroded depth as:

$$\begin{aligned}\dot{e}(t) &= (1 - \alpha_{NE})\rho_{bed} \int_0^{L_{bed}} \frac{\partial H(x, t)}{\partial t} dx \\ &= (1 - \alpha_{NE})L_{bed}\rho_{bed} \frac{d\bar{H}}{dt},\end{aligned}\quad (21)$$

where  $L_{bed}$  is the length of the particle bed.

The emission rate is estimated by the weighing measurements of sand as:

$$\dot{e}(t) = \frac{\Delta M}{W_{bed} \cdot \Delta t}, \quad (22)$$

where  $\Delta M$  is the emitted mass during the time interval  $\Delta t$ .

Finally, the pavement phenomenon was also explored by means of digital analysis of the final pictures from the sand bed using a thresholding method and the commercial software Matlab. It was possible to visualise the accumulation of coarse black particles and to calculate the final cover rate.

## 4. Experimental results

### 1. Temporal evolution of the pavement process

Fig. 8 presents the influence of  $\alpha_{NE}$  and  $U_\infty$  on the emission rate obtained using Equation 22. During the first minutes, it remains nearly constant but decreases afterwards due to the pavement phenomenon. The constant flux during the first minutes can be explained using an analogy to steady state saltation, in which an equilibrium occurs between wind flow and transport. For the case with  $U_\infty = 9.6 \text{ m/s}$  and  $\alpha_{NE} = 20\%$  for the low velocity  $6.7 \text{ m/s}$ . However, for  $\alpha_{NE} = 20\%$  the phenomenon was clearly seen for all tested velocities.

**Figure 8.** Temporal evolution of the emitted mass flux for non-erodible mass fractions of 10 and 20%

From the experimental data of Fig. 8, two interesting features can be drawn: the duration  $T$  during which the emission rate  $\dot{e}(t)$  measured at the downwind end of the bed is constant and the characteristic time  $\tau$  which corresponds to the transient time characterising the decrease of the emission rate  $\dot{e}(t)$  to zero. This temporal variation is associated with the spatial variation of saltation, namely with the so-called saturation length.

**Figure 9.** Sketch of the successive positions of the erosion front that propagates from left to right (red lines). For simplicity, the erosion front is assumed to have a rectilinear shape.

The temporal scenario of the bed erosion process can be described as follows. The pavement phenomenon first appears at the upwind edge of the bed and progresses downwind over time. This erosion front extends over a finite length  $L_{front}$  due to the saturation process of saltation.  $L_{front}$  is thus expected to be proportional to the saturation length  $l_{sat}$ , typically two or three times  $l_{sat}$ . For definitiveness, we will set  $L_{front} = a l_{sat}$  with  $a = 3$ . Upwind the front, the erosion depth is  $z = H_f$  and downwind, it is  $z = 0$ . In between, the bed surface evolves continuously from  $z = H_f$  to  $z = 0$  (see Fig. 9).

The position of the downwind edge of the erosion front will be denoted by  $x_{down}$  (The upwind

edge is thus  $x_{up} = x_{down} - L_{front}$ .

As soon as the downwind edge of the erosion front has not reached the downwind end of the bed, the emission rate is equal to the saturation flux  $Q_{sat}$ :

$$\dot{e}(t) = Q_{sat} , \quad (23)$$

The time  $T$  needed from the downwind edge of the erosion front to reach the downwind end of the bed is simply given by the total mass of sediment eroded over the distance  $L_{bed} - L_{front}/2$  divided by the saturation flux:

$$T = \frac{(1 - \alpha)\rho_{bed}H_f(L_{bed} - L_{front}/3)}{Q_{sat}} . \quad (24)$$

Let assume for simplicity that the erosion front has a linear shape:  $h(x) = H_f x/L_{front}$  with  $x = x_{up} - c t$ , where  $c$  is the front propagation velocity. From mass conservation, we easily obtain that the variation of the mass flux  $Q(x)$  along the erosion front is given by:

$$Q(x) = \frac{x}{L_{front}} Q_{sat} . \quad (25)$$

and the front celerity  $c$  by:

$$c = \frac{Q_{sat}}{(1 - \alpha)\rho_{bed}H_f} . \quad (26)$$

The emission rate  $\dot{e}(t)$  is expected to decrease when the downwind edge of the erosion front has reached the end of the bed. If we assume that Equation 25 still holds as the erosion front approaches the end of the bed,  $\dot{e}(t)$  corresponds to the value taken by the flux  $Q(x)$  at  $x = L_{front} - c(t - T)$  which yields:

$$\dot{e}(t) = Q_{sat} \left(1 - \frac{t-T}{\tau}\right) , \quad (27) \text{ with: } \tau = \frac{(1 - \alpha)\rho_{bed}H_f L_{front}}{Q_{sat}} , \quad (28)$$

for  $T < t < T + \tau$ .

Therefore,  $\dot{e}(t)$  decreases linearly with time from  $t = T$ .  $\tau$  represents the characteristic time needed for the emission rate  $\dot{e}$  to vanish. The experimental data seems to indicate that the decrease of  $\dot{e}(t)$  is better described by an exponential decrease rather than a linear one. A more elaborate calculation would probably provide a prediction in better agreement with the experiments. However, this crude calculation gives insights on how the characteristic times  $T$  and  $\tau$  are related to the physical parameters of the problem. In particular, the knowledge of these characteristic times allows to provide an estimate for the saturation length  $l_{sat}$  and the final erosion depth  $H_f$ :

$$l_{sat} = \frac{L_{bed}}{a(0.5 + T/\tau)} , \quad (29)$$

$$H_f = \frac{(0.5\tau + T)Q_{sat}}{(1 - \alpha)\rho_{bed}L_{bed}} . \quad (30)$$

The values of  $l_{sat}$  and  $H_f$  deduced from the above relationships are reported in Table 2 in comparison with the experimental values of the final eroded depth  $H_f$  from the laser measurements. This provides fairly reasonable values.

**Table 2.** Values of the saturation length  $l_{sat}$  and  $H_f$  deduced from Equations 29-30 for the different sets of experiments.

$\alpha_{NE}$	10%			20%		
U (m/s)	6.7	8.5	9.6	6.7	8.5	9.6
T (min)	4.5	4.5	3.9	4.5	2.5	1.0
$Q_{sat}$ (kg/min.m)	0.94	2.45	4.9	1.09	2.43	5.04
$\tau$ (min)	9.8	3.3	1.7	2.2	1.4	1.4
$H_f$ (mm)	1.03	1.75	2.5	0.88	1.19	1.39
(Experimental)						
$l_{sat}$ (m)	2.3	1.2	0.8	0.87	0.98	1.8
$H_f$ (mm)	0.93	1.59	2.25	0.73	0.93	1.02
(Equation 30)						

Fig. 10 presents the temporal evolution of the eroded depths of the sand bed obtained by laser measurements for two different cases. Notice that the erosion process is observed in the last part of the bed (between  $5.50 \text{ m} < x < 5.82 \text{ m}$ ). Fig. 10(a) shows a sharp drop in the eroded depth ( $H$ ) at 4 minutes and the pavement after 5 minutes. Fig. 10(b) shows similar behaviour but at different times. The same behaviour was observed in all tested cases, more or less sharply depending on the velocity and on  $\alpha_{NE}$ . These results enable a comprehension of how pavement occurs and are consistent with the results shown in Fig. 8. During the first minutes of the experiment, there were emissions along the whole tunnel. The eroded depth in the test section did not have significant changes since depositions also occurred. As the pavement phenomenon develops, initially at the beginning of the tunnel, emissions from the earlier portions decrease, tending to zero. At these moments, there would be no more deposition in the test section. Therefore, there is a sudden drop in the eroded depth, which corresponds to the stage when the pavement occurs at the part of the sand bed located in the test section (close to the end of the tunnel). As a consequence of the spatial differences on the temporal evolution of the eroded depth, it is difficult to estimate the emission rate using Equation 21, considering that there are laser measurements of a small part of the bed (320 mm, in the test section). Nonetheless, assuming that the final eroded depth is spatially constant, the emitted mass  $E_f$  can be estimated from laser measurements. Table 3 shows the experimental results of  $H_f$  and the results of  $E_f$  obtained from  $H_f$  measurements and by weighing. The different procedures provide close values of  $E_f$ , demonstrating good measurement accuracy.

**Figure 10.** Temporal evolution of the eroded depths for two different cases: (a)  $U_{\infty} = 9.6 \text{ m/s}$  with  $\alpha_{NE} = 10\%$  and (b)  $U_{\infty} = 8.5 \text{ m/s}$  with  $\alpha_{NE} = 20\%$

**Table 3.** Final eroded depths of the sand bed determined by laser methodology and estimated emitted mass obtained by  $H_f$  measurements and by weighing

$\alpha_{NE}(\%)$	$U_{\infty}(\text{m/s})$	$H_f(\text{mm})$	$E_f(\text{g})$ (from weighing)	$E_f(\text{g})$ (from $H_f$ )
10	6.7	1.03	2401.3	2393.3
	8.5	1.75	4074.8	4140.9
	9.6	2.5	5821.2	6095.2

20	6.7	0.88	1821.4	1719.0
	8.5	1.19	2461.0	2505.7
	9.6	1.34	2773.5	3088.7

## 2. Final state of the particles bed after the pavement phenomenon

Fig. 11 presents pictures of the sand bed before and after the pavement for the simulated case with  $U_{\infty} = 8.5 \text{ m/s}$  and  $\alpha_{NE} = 20\%$ . Digital analysis of these photographs provided data of final cover rates. The results are presented in Table 4.

**Figure 11.** Pictures of the sand bed (a) before and (b) after the pavement for  $U_{\infty} = 8.5 \text{ m/s}$  with  $\alpha_{NE} = 20\%$

Previous authors such as [12] have assumed that erosion ceases when the bed reaches a final cover rate independently of velocity and  $\alpha_{NE}$ . However, the results presented in Table 4 show that  $CR_f$  varies with  $\alpha_{NE}$ . In addition, the final cover rates and eroded depths of the sand bed increase with velocity, as greater velocities cause more emissions and consequently greater accumulation of non-erodible particles.

For  $\alpha_{NE} = 10\%$ , the values of  $CR_f$  were lower than those found for  $\alpha_{NE} = 20\%$ . Nonetheless, the eroded depths are greater for  $\alpha_{NE} = 10\%$ . These results support the fact that the protective role of the non-erodible particles is not only given by the cover rate of the sand bed but also by the height of the non-erodible elements (see Equation 7).

**Table 4.** Initial and final cover rates determined by digital analysis of photographs of the experiment and final eroded depths of the sand bed as determined by laser methodology ( $H_f$ )

$\alpha_{NE}(\%)$	$CR_i(\%)$	$U_{\infty}(\text{m/s})$	$CR_f(\%)$	$H_f(\text{mm})$
10	6.4	6.7	17	1.03
		8.5	23	1.75
		9.6	30	2.50
20	11.8	6.7	29	0.88
		8.5	35	1.19
		9.6	38	1.34

Fig. 12 shows the ratio between  $CR_f$  and  $CR_i$  for each tested case. It can be seen that despite the lower values of  $CR_f$  for  $\alpha_{NE} = 10\%$ , the increase in cover rate was proportionally more steep. Moreover, the airflow velocity has more influence on the cover rate variation for the case with  $\alpha_{NE} = 10\%$ .

**Figure 12.** Variation of the cover rate for each tested configuration

## 5. Comparisons between model predictions and experiments

As explained in Section 3, the emitted mass is estimated by following three steps. (i) First, knowing the initial particle size distribution, the cover rate  $CR$  is written as a function of the eroded depth  $H$  (Equation 17). Fig. 13 presents  $CR/CR_i$  as a function of  $H$  obtained using Equation 17 (solid line) and the experimental data, which is the final state of the bed ( $H_f$  and  $CR_f/CR_i$ ) for each tested case (points). The formulation given by Equation 17 is well supported by the experiments. (ii) Then, Equation 17 enables the calculation of the final eroded depth after the pavement phenomenon for a given velocity using Equation 19 (in the experimental configuration,  $u_{MIN}^* = u_{td}^*(200\mu\text{m}) \approx 0.186 \text{ m/s}$ ). It is important to state that

both equations produce differences between the modelled and experimental results. As a consequence, the propagated error of the final eroded depth leads to a variation in the emission estimation. (iii) Finally, emission is calculated using Equation 20.

**Figure 13.** Cover rate of non-erodible particles as a function of eroded depth obtained by modelling (Equation 17) and the values of  $H_f$  and  $CR_f/CR_i$  obtained experimentally (Table 4)

Table 5 compares the estimated results of the final eroded depth (Equation 19) and the emitted mass with experimental data, including percentage errors. There is good agreement between the modelled and the experimental results for both measurements. The uncertainty of emission measurements, retrieved by previous repeatability tests with similar wind tunnel experiments, was found to be in the range of  $\pm 6.5\%$ . Thus, the modelled results slightly underestimate the experimental data.

**Table 5. Comparison of experimental and model results**

$\alpha_{NE}$	$U_\infty$	$H_f$ (mm)			Emitted mass (g)		
		(m/Exp.	Mod.	Error	Exp.	Mod.	Error
10%	6.7	1.03	0.98	4.9%	2393.32	345.1	2.0%
	8.5	1.75	1.73	1.1%	4140.93	802.9	8.2%
	9.6	2.50	2.23	10.8%	6095.25	231.8	14.2%
20%	6.7	0.88	0.50	43.2%	1719.01	049.5	38.9%
	8.5	1.19	0.99	16.8%	2505.72	092.5	16.5%
	9.6	1.34	1.26	6.0%	3088.72	637.3	14.6%

The proposed model is a continuation of the work described in [12]. A limitation of the previous model is the assumption that erosion is finalised when the bed is completely overlaid by non-erodible particles. As explained in Section 2, it has already been verified that erosion is terminated with a lower cover rate of the surface. Another limitation of the previous [12] model is related to the fact that the influence of saltation on transport is not accounted for. These two points have been adjusted in [12], but with a purely stochastic method that produces very large calculation times. On the other hand, the previous model has the advantage of providing detailed information concerning the temporal decrease of the emitted mass flux, although in practical cases such as fugitive sources on industrial sites, the most important information is the total emitted mass. Furthermore, the absolute value of the emitted mass is largely underestimated by the [12] model. The present study focused on modelling the final state of the bed after the pavement phenomenon and provides estimates of the global emissions, which are in good agreement with the experimental results. Moreover, it has the advantage of being represented by a simple mathematical expression.

## 6. Conclusion

This work investigated the pavement phenomenon and how it affects particle emissions from a bed of granular material exposed to a turbulent flow using experimental and analytical approaches. Wind erosion can lead to serious environmental impacts since the emissions can extend for large distance from the source. Literature review showed that dust emissions from diffuse sources are significant and may lead to air pollution episodes and consequently human healthy deterioration. Therefore, it is essential to have a proper emission model based on the physics of the phenomena and validated by experimental measurements.

The temporal evolution of the emitted mass flux revealed that during the first minutes, flux remained nearly constant, indicating that an equilibrium was established between wind flow



and transport. The equilibrium state lasted longer for the lower velocities and for the beds with a lower proportion of non-erodible particles. After a few minutes, the flux began to decrease due to the pavement phenomenon. This decay was faster for higher velocities. In addition, greater amounts of non-erodible particles in the mixture led to a greater decay rate. These findings are consistent with previous studies in the literature [11,16].

A mathematical model for emitted mass estimates that includes the influence of the presence of non-erodible particles and saltation was proposed. The influence of saltation was included in the modelling using the dynamic threshold. The proportion of non-erodible particles on a bed surface after the pavement phenomenon was quantified. Although previous authors (e.g., [11,16]) have observed this phenomenon, it had not yet been quantified. Good agreement was found between the experimental and model results for the global emissions and bed eroded depth. Although the theory on which the proposed model is based is quite complex, the model is represented by a simple algebraic expression and demands low computational effort. Further work is still needed to improve the model for application to stockpiles of granular materials.

## Acknowledgements

The authors acknowledge support from Capes/Cofecub, Fapes and CNPq.

## References

- [1] Alfaro, C., Gomes, L., 2001. Modeling mineral aerosol production by wind erosion : Emission intensities and aerosol size distributions in source areas. *Journal of Geophysical Research* 106, 75–84.
- [2] Avecilla, F., Panebianco, J.E., Mendez, M.J., Buschiazzo, D.E., 2018. PM10 emission efficiency for agricultural soils: Comparing a wind tunnel, a dust generator, and the open-air plot. *Aeolian Research* 32, 116–123. doi:10.1016/j.aeolia.2018.02.003.
- [3] Bagnold, R.A., 1941. The physics of blown sand and desert dunes, volume 265. Reprint ed., Methuen, London.
- [4] Brown, S., Nickling, W.G., Gillies, J.A., 2008. A wind tunnel examination of shear stress partitioning for an assortment of surface roughness distributions. *Journal of Geophysical Research: Earth Surface* 113, 1–13. doi:10.1029/2007JF000790.
- [5] Carvacho, O.F., Ashbaugh, L.L., Brown, M.S., Flocchini, R.G., 2004. Measurement of PM2.5 emission potential from soil using the UC Davis resuspension test chamber. *Geomorphology* 59, 75–80. doi:10.1016/j.geomorph.2003.09.007.
- [6] Charru, F., Andreotti, B., Claudin, P., 2013. Sand Ripples and Dunes. *Annual Review of Fluid Mechanics* 45, 469–493. doi:10.1146/annurev-fluid-011212-140806.
- [7] Chepil, W.S., 1945. Dynamics of wind erosion: I. Nature of movement of soil by wind. *Soil Science* 60, 305–320. doi:00010694-194510000-00004.
- [8] Claudin, P., Andreotti, B., 2006. A scaling law for aeolian dunes on Mars, Venus, Earth, and for subaqueous ripples. *Earth and Planetary Science Letters* 252, 30–44. doi:10.1016/j.epsl.2006.09.004, arXiv:0603656.
- [9] Courrech du Pont, S., 2015. Dune morphodynamics. *Comptes Rendus Physique* 16, 118–138. doi:10.1016/j.crhy.2015.02.002.
- [10] Crawley, D.M., Nickling, W.G., 2003. Drag partition for regularly-arrayed rough surfaces. *Boundary-Layer Meteorology* 107, 445–468. doi:10.1023/A:1022119909546, arXiv:arXiv:1011.1669v3.
- [11] Descamps, I., 2004. Erosion éolienne d'un lit de particules à large spectre granulométrique. Wind erosion of a bed of particles with a wide size distribution. Ph.D. thesis. Ecole des Mines de Douai, Université de Valenciennes et du Hainaut-Cambrésis.
- [12] Descamps, I., Harion, J.L., Baudoin, B., 2005. Taking-off model of particles with a wide size distribution . *Chemical Engineering and Processing: Process Intensification* 44, 159–166. doi:10.1016/j.cep.2004.04.007.
- [13] Duan, S., Cheng, N., Xie, L., 2013. A new statistical model for threshold friction velocity of sand particle motion. *CATENA* 104, 32 – 38.
- [14] Durán, O., Claudin, P., Andreotti, B., 2011. On aeolian transport: Grain-scale interactions, dynamical mechanisms and scaling laws. *Aeolian Research* 3, 243 – 270. doi:10.1016/j.aeolia.2011.07.006.
- [15] Foucaut, J.M., Stanislas, M., 1996. Take-off threshold velocity of solid particles lying under a turbulent boundary layer. *Experiments in Fluids* 20, 377–382. doi:10.1007/BF00191019.
- [16] Furieri, B., Harion, J., Milliez, M., Russeil, S., Santos, J., 2014a. Numerical modelling of aeolian erosion over a surface with non-uniformly distributed roughness elements. *Earth Surface Processes and Landforms* 39. doi:10.1002/esp.3435.
- [17] Furieri, B., Russeil, S., Santos, J., Harion, J., 2013. Effects of non-erodible particles on aeolian erosion: Wind-tunnel simulations of a sand oblong storage pile. *Atmospheric Environment* 79, 672–680. doi:10.1016/j.atmosenv.2013.07.026.
- [18] Furieri, B., Santos, J.M., Russeil, S., Harion, J.L., 2014b. Aeolian erosion of storage

- piles yards: Contribution of the surrounding areas.
- Environmental Fluid Mechanics*
- 14, 51–67. doi:10.1007/s10652-013-9293-4.
- [19] Gillette, D.A., Stockton, P.H., 1989. The effect of nonerodible particles on wind erosion of erodible surfaces. *Journal of Geophysical Research: Atmospheres* 94, 12885–12893. doi:10.1029/JD094iD10p12885.
- [20] Hamdan, M.A., Refaat, A.A., Abdel Wahed, M., 2016. Morphologic characteristics and migration rate assessment of barchan dunes in the Southeastern Western Desert of Egypt. *Geomorphology* 257, 57–74. doi:10.1016/j.geomorph.2015.12.026.
- [21] Ho, T.D., Valance, A., Dupont, P., Ould El Moctar, A., 2014. Aeolian sand transport: Length and height distributions of saltation trajectories. *Aeolian Research* 12, 65–74. doi:10.1016/j.aeolia.2013.11.004.
- [22] Iversen, J.D., Rasmussen, K.R., 1994. The effect of surface slope on saltation threshold. *Sedimentology* 41, 721–728. doi:10.1111/j.1365-3091.1994.tb01419.x.
- [23] Iversen, J.D., Wang, W.P., Rasmussen, K.R., Mikkelsen, H.E., Leach, R.N., 1991. Roughness element effect on local and universal saltation transport, in: Barndorff-Nielsen, O., Willetts, B. (Eds.), *Aeolian Grain Transport*. Springer Vienna. volume 2 of *Acta Mechanica Supplementum*, pp. 65–75. doi:10.1007/978-3-7091-6703-8\_5.
- [24] Iversen, J.D., White, B.R., 1982. Saltation threshold on Earth, Mars and Venus. *Sedimentology* 29, 111–119. doi:10.1111/j.1365-3091.1982.tb01713.x.
- [25] Kheirabadi, H., Mahmoodabadi, M., Jalali, V., Naghavi, H., 2018. Sediment flux, wind erosion and net erosion influenced by soil bed length, wind velocity and aggregate size distribution. *Geoderma* 323, 22–30. doi:10.1016/j.geoderma.2018.02.042.
- [26] Kok, J., Renno, N., 2006. Enhancement of the emission of mineral dust aerosols by electric forces. *Geophysical research letters* 33, 1–5. doi:10.1029/2006GL026284.
- [27] Kok, J.F., Parteli, E.J.R., Michaels, T.I., Karam, D.B., 2012. The physics of wind-blown sand and dust. *Reports on Progress in Physics* 75, 106901. doi:10.1088/0034-4885/75/10/106901.
- [28] Kok, J.F., Renno, N.O., 2009. A comprehensive numerical model of steady state saltation (COMSALT). *Journal of Geophysical Research: Atmospheres* 114, 1–20. doi:10.1029/2009JD011702.
- [29] Li, B., McKenna Neuman, C., 2014. A wind tunnel study of aeolian sediment transport response to unsteady winds. *Geomorphology* 214, 261–269. doi:10.1016/j.geomorph.2014.02.010.
- [30] Li, L., Martz, L.W., 1995. Aerodynamic dislodgement of multiple-size sand grains over time. *Sedimentology* 42, 683–694. doi:10.1111/j.1365-3091.1995.tb00400.x.
- [31] Lima, I.A., Araújo, A.D., Parteli, E.J., Andrade, J.S., Herrmann, H.J., 2017. Optimal array of sand fences. *Scientific Reports* 7, 1–8. doi:10.1038/srep45148.
- [32] Liu, L.Y., Shi, P.J., Zou, X.Y., Gao, S.Y., Erdon, H., Yan, P., Li, X.Y., Dong, Z.B., Wang, J.H., 2003. Short-term dynamics of wind erosion of three newly cultivated grassland soils in Northern China. *Geoderma* 115, 55 – 64. doi:10.1016/S0016-7061(03)00075-2.
- [33] Lyles, L., Schrandt, R.L., Schmeidler, N.F., 1974. How aerodynamic roughness elements control sand movement. *Transactions of the American Society for Agricultural Engineers* 17, 134–139. doi:10.13031/2013.36805.
- [34] Marshall, J.K., 1971. Drag measurements in roughness arrays of varying density and distribution. *Agricultural Meteorology* 8, 269–292. doi:10.1016/0002-1571(71)90116-6.
- [35] Marticorena, B., Bergametti, G., 1995. Modeling the atmospheric dust cycle: 1. Design of a soil-derived dust emission scheme. *Journal of Geophysical Research: Atmospheres* 100, 16415–16430. doi:10.1029/95JD00690.
- [36] Martin, R.L., Kok, J.F., Hugenholtz, C.H., Barchyn, T.E., Chamecki, M., Ellis, J.T., 2018. High-frequency measurements of aeolian saltation flux: Field-based methodology and

applications. *Aeolian Research* 30, 97–114. doi:10.1016/J.AEOLIA.2017.12.003.

[37] Mayaud, J.R., Bailey, R.M., Wiggs, G.F., Weaver, C.M., 2017. Modelling aeolian sand transport using a dynamic mass balancing approach. *Geomorphology* 280, 108–121. doi:10.1016/j.geomorph.2016.12.006.

[38] McKenna Neuman, C., Sanderson, R.S., Sutton, S., 2013. Vortex shedding and morphodynamic response of bed surfaces containing non-erodible roughness elements. *Geomorphology* 198, 45–56. doi:10.1016/j.geomorph.2013.05.011.

[39] Musick, H.B., Trujillo, S.M., Truman, C.R., 1996. Wind-tunnel modelling of the influence of vegetation structure on saltation threshold. *Earth Surface Processes and Landforms* 21, 589–605. doi:10.1002/(SICI)1096-9837(199607)21:7<589::AID-ESP659>3.0.CO;2-1.

[40] Neuman, C.M., Bédard, O., 2015. *Journal of Geophysical Research : Earth Surface* , 1–17doi:10.1002/2015JF003475.Received.

[41] Nickling, W.G., Neuman, C.M., 2009. Aeolian sediment transport, in: Parsons, A., Abrahams, A. (Eds.), *Geomorphology of Desert Environments*. Springer, pp. 517–555. doi:10.1007/978-1-4020-5719-9\_17.

[42] Raffaele, L., Bruno, L., Wiggs, G.F., 2018. Uncertainty propagation in aeolian processes: From threshold shear velocity to sand transport rate. *Geomorphology* 301, 28–38. doi:10.1016/j.geomorph.2017.10.028.

[43] Raupach, M.R., 1992. Drag and drag partition on rough surfaces. *Boundary-Layer Meteorology* 60, 375–395. doi:10.1007/BF00155203.

[44] Raupach, M.R., Gillette, D.A., Leys, J.F., 1993. The effect of roughness elements on wind erosion threshold. *Journal of Geophysical Research: Atmospheres* 98, 3023–3029. doi:10.1029/92JD01922.

[45] Schlichting, H., 1936. Experimentelle Untersuchungen zum Rauheitsproblem. *Ingenieur-Archiv* 7, 1–34. doi:10.1007/BF02084166.

[46] Seinfeld, J.H., Pandis, S.N., 2006. *Atmospheric Chemistry and Physics: from Air Pollution to Climate Change*. volume 51. Second ed., John Wiley & Sons, Inc. doi:10.1016/0016-7037(87)90252-3.

[47] Shao, Y., Lu, H., Shag, Y., 2000. A simple expression for wind erosion threshold friction velocity. *Journal of Geophysical Research* 105, 22437–22443. doi:https://doi.org/10.1029/2000JD900304.

[48] Smith, I.B., Spiga, A., Holt, J.W., 2015. Aeolian processes as drivers of landform evolution at the South Pole of Mars. *Geomorphology* 240, 54–69. doi:10.1016/j.geomorph.2014.08.026.

[49] Swet, N., Katra, I., 2016. Reduction in soil aggregation in response to dust emission processes. *Geomorphology* 268, 177–183. doi:10.1016/j.geomorph.2016.06.002.

[50] Turpin, C., Badr, T., Harion, J.L., 2010. Numerical modelling of aeolian erosion over rough surfaces. *Earth Surface Processes and Landforms* 35, 1418–1429. doi:10.1002/esp.1980.

[51] Turpin, C., Harion, J.L., 2009. Numerical modeling of flow structures over various flat-topped stockpiles height: Implications on dust emissions. *Atmospheric Environment* 43, 5579 – 5587. doi:10.1016/j.atmosenv.2009.07.047.

[52] Walter, B., Gromke, C., Lehning, M., 2012. Shear-Stress Partitioning in Live Plant Canopies and Modifications to Raupach's Model. *Boundary-Layer Meteorology* 144, 217–241. doi:10.1007/s10546-012-9719-4.

[53] Webb, N.P., Galloza, M.S., Zobeck, T.M., Herrick, J.E., 2016. Threshold wind velocity dynamics as a driver of aeolian sediment mass flux. *Aeolian Research* 20, 45–58. doi:10.1016/j.aeolia.2015.11.006.

[54] Webb, N.P., Strong, C.L., 2011. Soil erodibility dynamics and its representation for wind erosion and dust emission models. *Aeolian Research* 3, 165–179.

doi:10.1016/j.aeolia.2011.03.002.

[55] WHO, 2006. The World Health Report. Technical Report. World Health Organization.

[56] Wolfe, S.A., Nickling, W.G., 1996. Shear stress partitioning in sparsely vegetated desert canopies. *Earth Surface Processes and Landforms* 21, 607–619. doi:10.1002/(SICI)1096-9837(199607)21:7<607::AID-ESP660>3.0.CO;2- 1.

[57] Yang, F., Yang, X.H., Huo, W., Ali, M., Zheng, X.Q., Zhou, C.L., He, Q., 2017. A continuously weighing, high frequency sand trap: Wind tunnel and field evaluations. *Geomorphology* 293, 84–92. doi:10.1016/j.geomorph.2017.04.008.

[58] Zobeck, T.M., Sterk, G., Funk, R., Rajot, J.L., Stout, J.E., Van Pelt, R.S., 2003. Measurement and data analysis methods for field-scale wind erosion studies and model validation. *Earth Surface Processes and Landforms* 28, 1163–1188. doi:10.1002/esp.1033.

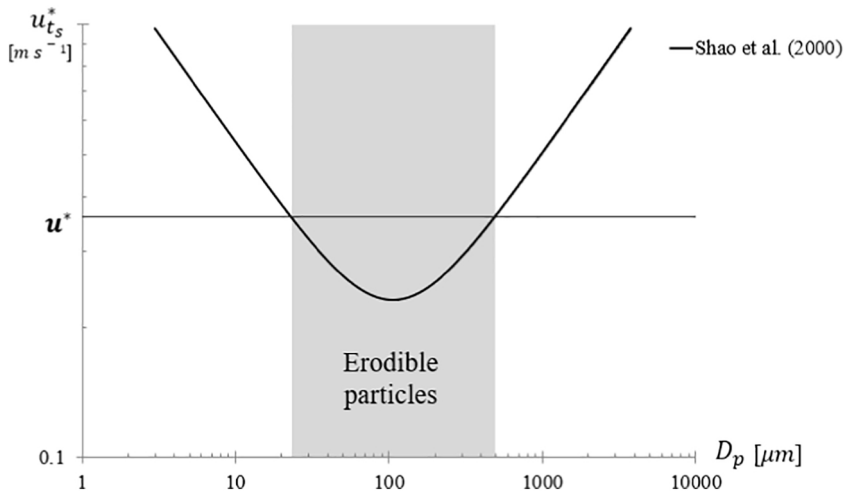


Figure 1

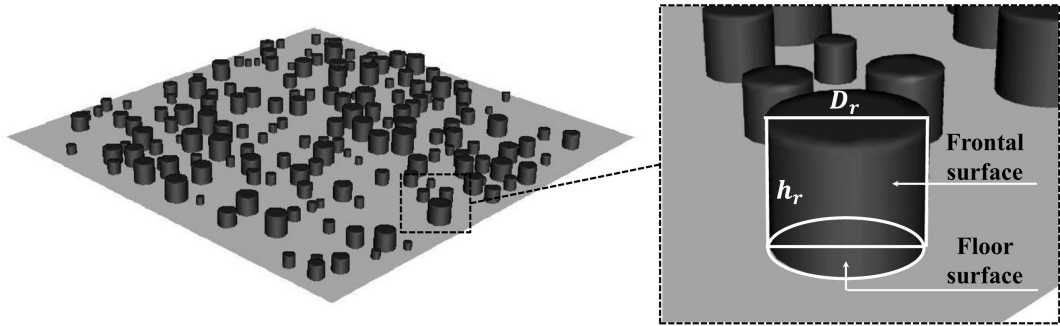


Figure 2

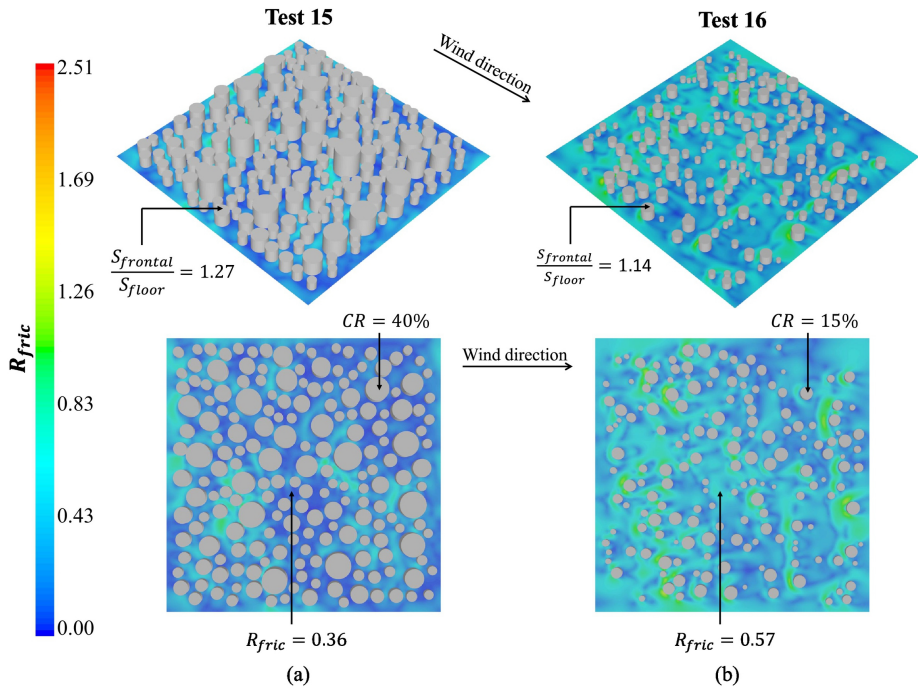


Figure 3



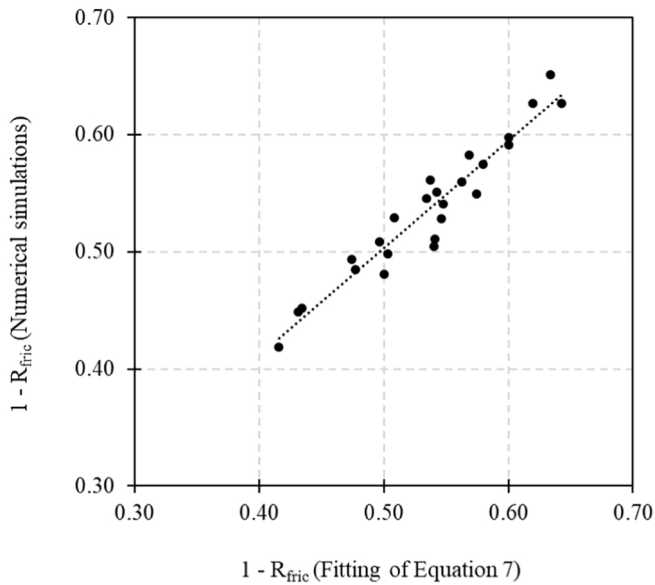


Figure 4

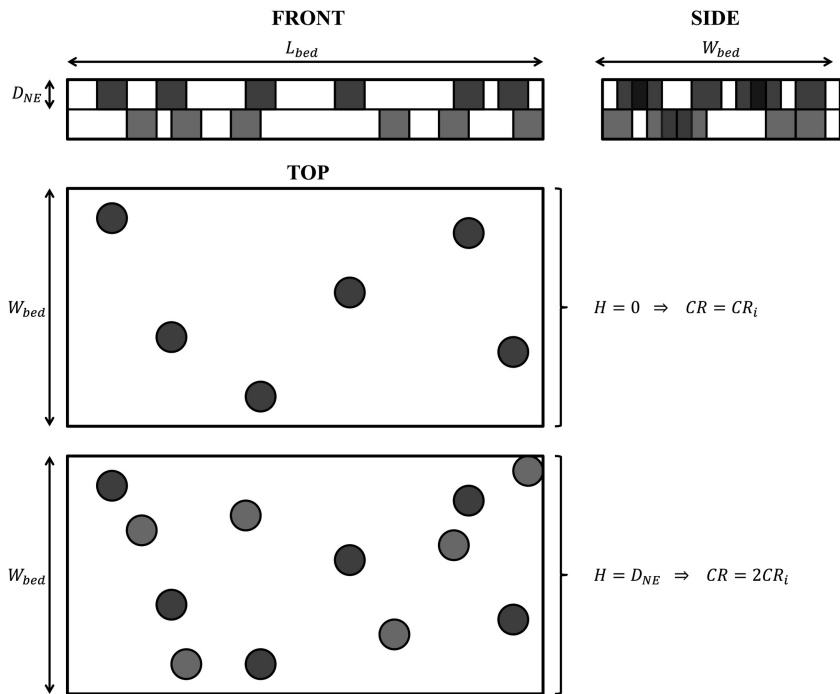


Figure 5

## INPUT VARIABLES

- Wind velocity
- Particle size distribution

- Mass fraction percentages  $\alpha_i$
- Volume fraction  $\varphi$
- Bed dimensions

Erodibility of the particles  
(Eq. 2)

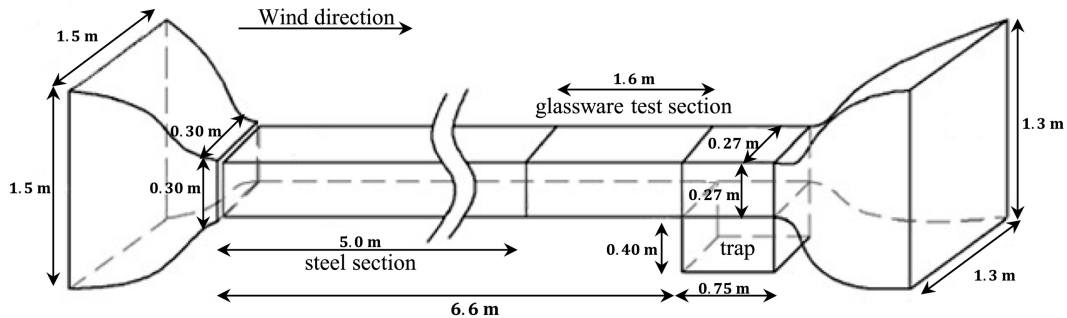
Distribution of the particles  
over the bed height (Eq. 15)

Relation between the eroded depth  
and the cover rate of NEP (Eq. 17)

Calculate  $H_f$  (Eq. 19)

Calculate  $E_f$  (Eq. 20)

Figure 6



(a)



(b)



(c)



(d)

Figure 7

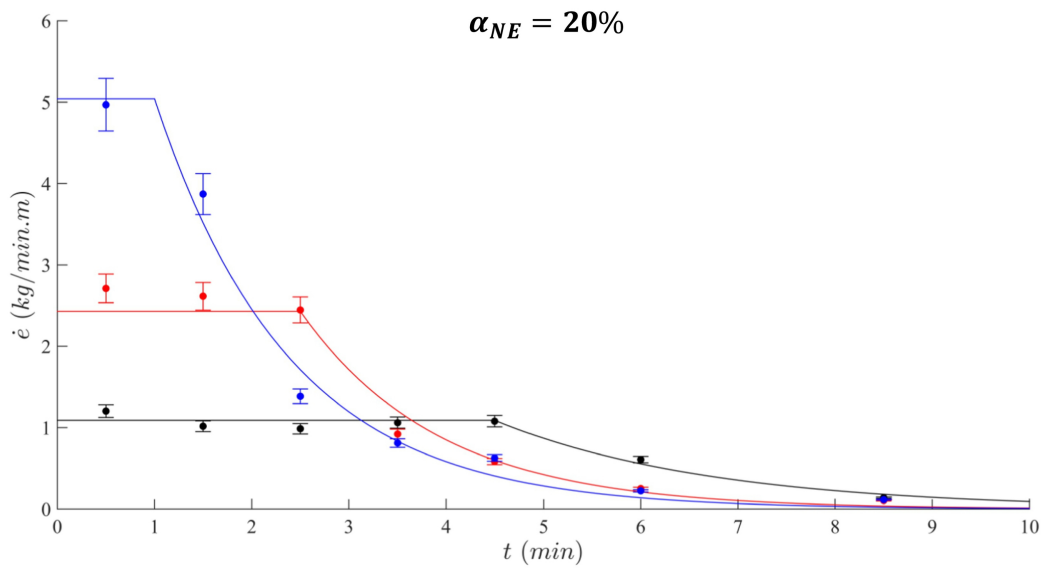
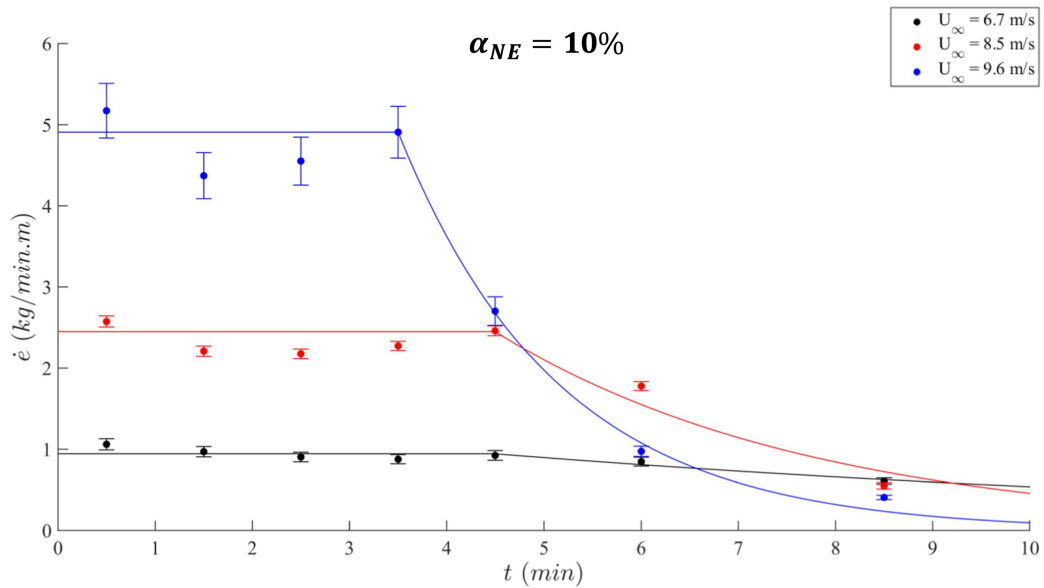


Figure 8

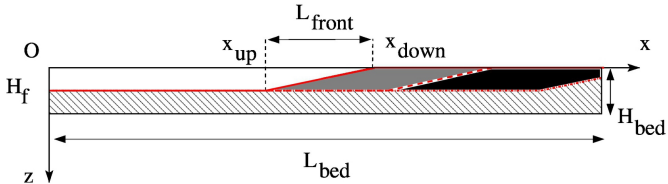
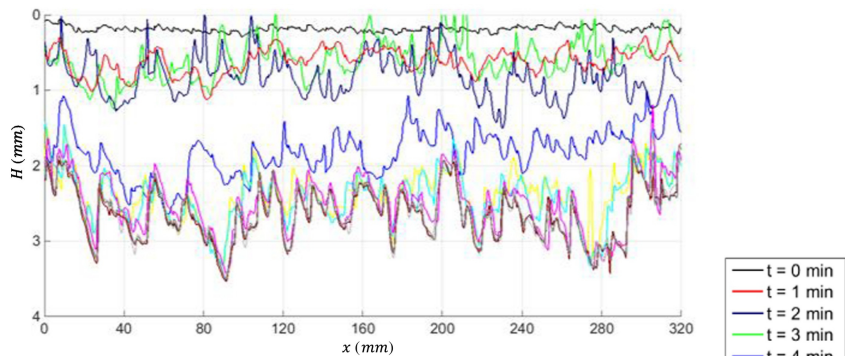
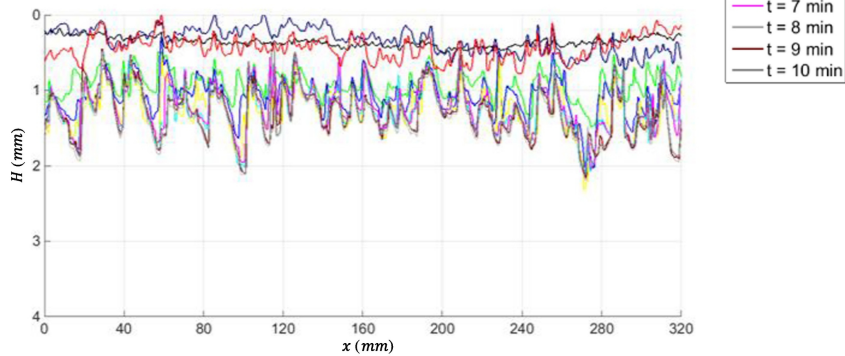


Figure 9

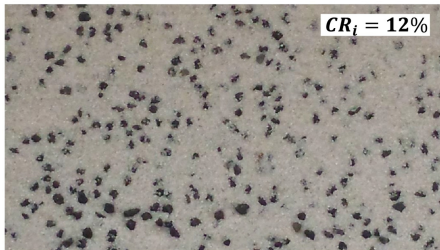


(a)



(b)

Figure 10



(a)



(b)

Figure 11



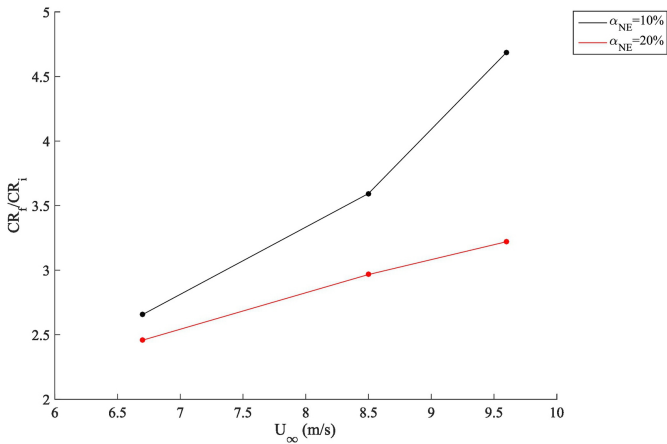


Figure 12

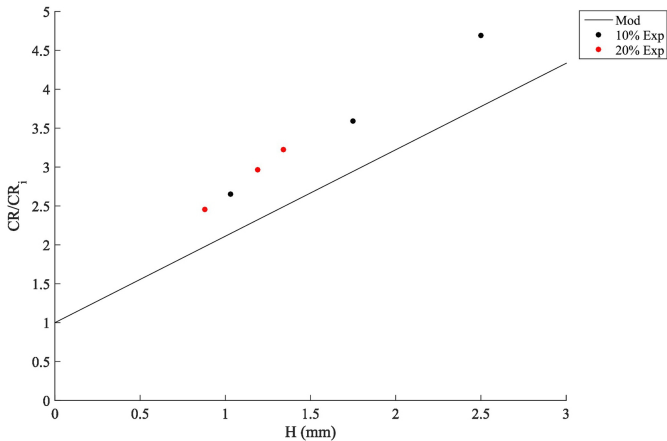


Figure 13